

## Present Day Kinematics of the Eastern California Shear Zone from a Geodetically Constrained Block Model

S. C. McClusky<sup>1</sup>, S. C. Bjornstad<sup>2</sup>, B. H. Hager<sup>1</sup>, R. W. King<sup>1</sup>, B. J. Meade<sup>1</sup>,  
M. M. Miller<sup>3</sup>, F. C. Monastero<sup>2</sup>, and B. J. Souter<sup>4</sup>

**Abstract.** We use Global Positioning System (GPS) data from 1993–2000 to determine horizontal velocities of 65 stations in eastern California and western Nevada between 35° and 37° N. We relate the geodetic velocities to fault slip rates using a block model that enforces path integral constraints over geologic and geodetic time scales and that includes the effects of elastic strain accumulation on faults locked to a depth of 15 km. The velocity of the Sierra Nevada block with respect to Nevada is  $11.1 \pm 0.3$  mm/yr, with slip partitioned across the Death Valley, ( $2.8 \pm 0.5$  mm/yr), Panamint Valley ( $2.5 \pm 0.8$  mm/yr), and Airport Lake/Owens Valley ( $5.3 \pm 0.7/4.6 \pm 0.5$  mm/yr) faults. The western Mojave block rotates at  $2.1 \pm 0.8^\circ/\text{My}$  clockwise, with  $3.7 \pm 0.7$  mm/yr of left lateral motion across the western Garlock Fault. We infer  $11 \pm 2$  mm/yr of right lateral motion across the Mojave region of the Eastern California Shear Zone.

### Introduction

About 25% of the relative motion between the Pacific and North American plates occurs east of the San Andreas fault zone (SAFZ) within a complex region characterized geologically [Dokka and Travis, 1990] and geodetically [Savage et al., 1990] as the Eastern California shear zone (ECSZ). South of the Garlock fault, shear within the eastern Mojave Desert occurs on right lateral faults that are nearly parallel to the direction of plate motion [Saubert et al., 1994, Miller et al., 2001]. North of the Garlock fault, in the southern Walker Lane belt, a series of faults accommodates both right lateral and normal motion. We present surface velocities estimated from GPS measurements of this region and relate these geodetic velocities to geologic fault slip rates using a block model.

### GPS Measurements and Analysis

The principal data used for this study were acquired in 14 surveys conducted from 1993–2000 by the National Imaging and Mapping Agency (NIMA) under the auspices of the Navy

Geothermal Program Office (GPO) and focused on the Coso region east of the Sierra Nevada Front. Each survey consisted of three to twelve 6-hour occupations of 23–34 monuments. Additional data come from annual surveys, 1993–1996 covering the broader Eastern California Shear Zone (ECSZ) region over the same time interval by Central Washington University (CWU) and the University of Miami (UM), [Miller et al., 2001; Dixon et al., 2000]. We also incorporated data from annual surveys, 1993–1996, in the Hunter Mountain, Panamint Valley and Death Valley regions by the Nuclear Regulatory Commission (NRC), California Institute of Technology, and the Harvard-Smithsonian Center for Astrophysics [Bennett et al., 1997] and along a transect measured by the United States Geological Survey (USGS) [Gan et al., 2000] in 1994, 1996, 1997, and 1999.

We analyze the data using the GAMIT/GLOBK software [King and Bock, 1999; Herring, 1999] in a two-step approach (e.g., Dong et al. [1998]). In the first step, we use GPS phase observations from each day to estimate station coordinates, the zenith delay of the atmosphere at each station, and orbital and Earth orientation parameters (EOP). In the second step we use the loosely constrained estimates of station coordinates, orbits, and EOP and their covariances from each day, aggregated by survey, as quasi-observations to estimate a consistent set of coordinates and velocities. For the analysis of all but the GPO data, we provide orbital control and tie the regional measurements to an external reference frame by including in the regional analysis data from 3–4 continuously operating stations of the International GPS Service (IGS) and by combining the regional quasi-observations with quasi-observations from an analysis of phase data from over 100 IGS stations performed by the Scripps Orbital and Permanent Array Center (SOPAC) at UC San Diego [Bock et al. 1997]. For the GPO surveys we do not mix the phase observations from the regional and IGS stations in order to avoid errors due to mismatched receivers and antennas. The most serious of these errors is an unknown but apparently constant offset in the sampling time of the codeless Ashtech LM-XII receivers used prior to November, 1996. This offset leads to position errors of 20–50 mm if the LM-XII data are differenced with those from other receivers, but has no significant effect on our estimates of relative positions within the GPO network. For our daily analysis of the GPO surveys we use IGS orbits constrained to one part in  $10^8$ . The GPO network is tied to an external reference frame in the second step of our analysis through 10 stations in common with the CWU/UM, NRC and/or USGS networks.

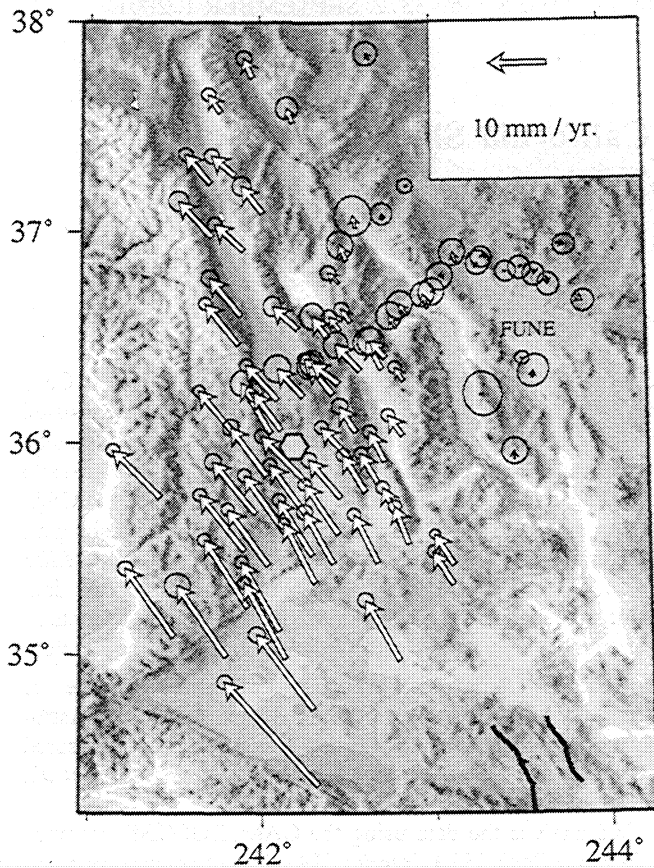
Before estimating velocities in the second step of our analysis, we examine the time series of position estimates to determine the appropriate weights to be applied to each group's surveys. For the velocity solution, we reweight the quasi-observations such that the normalized long-term scatter

<sup>1</sup> Department of Earth, Atmospheric, and Planetary Sciences, MIT, Cambridge, MA.

<sup>2</sup> Geothermal Program Office, Naval Air Weapons Station, China Lake, CA.

<sup>3</sup> Department of Geological Sciences, Central Washington University, Ellensburg, WA.

<sup>4</sup> Souter Computing Services, Baltimore, MD, 21239.



**Figure 1.** Horizontal velocity field for the ECSZ. Velocities are rotated into a North American reference frame and then shown relative to station FUNE. Heavy lines in the southeast (from west to east) show the locations of the 1992 Landers and 1999 Hector Mine Earthquake surface ruptures. The hexagonal symbol indicates the location of the Coso geothermal area.

in horizontal position for each group is unity. Finally, to account for correlated errors, we add to the assumed error in horizontal position a random walk component of  $0.5\text{mm}/\sqrt{\text{yr}}$ .

### Kinematic Model

To relate the geodetically determined velocities<sup>5</sup> (Figure 1) to geologic slip rates on faults, we must account for elastic strain accumulation. The classic model of a single, infinitely long fault [e.g., *Savage and Burford, 1973*] assumes that an elastic half-space is strained by slip imposed at the geologic rate on the down-dip extension of the fault below the locking depth,  $D$ . The geodetic velocity can be calculated as the difference between the geologic motions of the blocks and that caused by an equal slip deficit rate imposed on the fault above  $D$ . More realistic models incorporate a viscoelastic rheology characterized by a relaxation time,  $\tau$ , loaded by earthquakes with repeat time,  $T$ . Models of an elastic layer of thickness  $D$  and modulus  $\mu$  overlying a viscoelastic halfspace of viscosity  $\eta$ , with  $\tau = \eta/\mu$ , show little variation in interseismic velocities if  $\tau/T > 0.5$  [e.g., *Savage and Lisowski, 1998*], with the velocities closely approximating those of the classic model. If  $\tau/T < 0.2$ , significant variations arise that depend on  $\tau/T$  and the time,  $t$ , since the last earthquake. Near

the fault, early in the earthquake cycle, the geodetic rates are higher than the geologic rate, while late in the cycle, the apparent locking depth is much greater than  $D$ . In the far-field, the velocities approach the geologic rate.

We extend the *Savage and Burford (1973)* model to include the effects of multiple intersecting faults of finite length and varying orientations. As in their classic model, we assume that there is negligible variation with time of the interseismic velocity field. We use a block model [*Souter, 1998*, see *Meade et al., 2001*, for details] that requires that geologic surface displacements are the sum of coseismic and interseismic displacement fields, enforcing the same path integral constraint over both geodetic and geologic time scales. That is, the relative velocity between any two points does not depend on the path connecting them. Block boundaries are defined by faults with slip rate components determined in an internally consistent manner by the projection of the relative block velocity vector onto the fault plane. The surface deformation associated with backslip on these faults is modeled using dislocation theory [*Okada, 1985*], assuming a Poisson half-space. In general, the surface traces of the block boundaries are coarse representations of faults with Quaternary displacements [*Jennings, 1994*]. Individual slip rate estimates may reflect the total slip budget of several faults that are closely spaced. In our model all faults east of the SAF are vertical. *Souter [1998]* showed that the horizontal surface displacements calculated using fault-perpendicular displacements on a vertical dislocation are nearly the same as those calculated using a more realistic geometry with dip-slip displacements on dipping dislocations.

A velocity of  $\sim 2\text{mm/yr}$  associated with the SAFZ extends into the ECSZ region. We incorporate this effect by including surface velocity estimates from a model of California, west of the SAF [*Souter, 1998*; *B. J. Meade, unpublished results*], where block motions are estimated using the Southern California Earthquake Center (SCEC) horizontal deformation velocity map [*Shen, 1997*].

We estimate block motions using a weighted least squares inversion of the velocity vectors, with the data weighted by

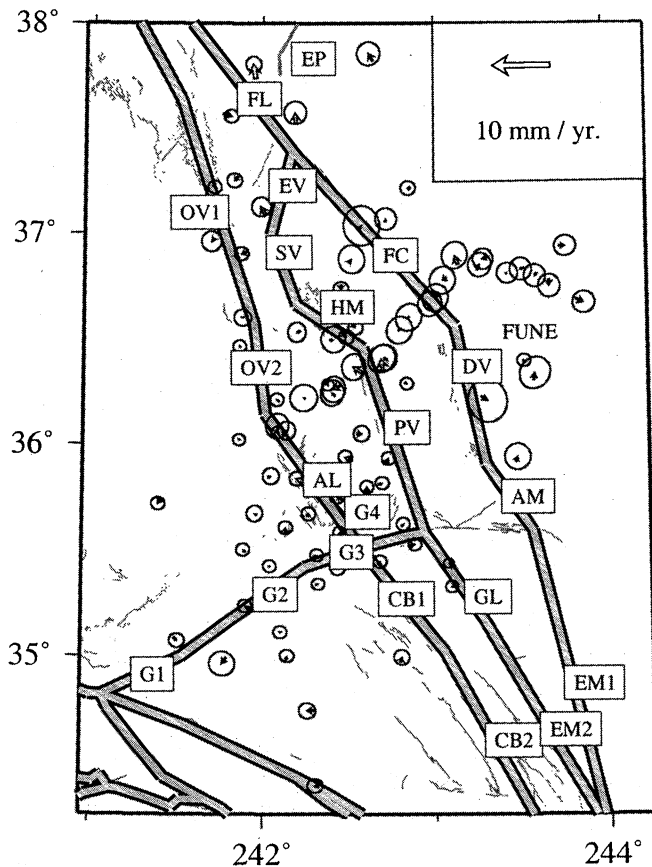
**Table 1.**

Fault name	Fault symbol	Strike slip rate <sup>a</sup> (mm/yr)	Fault normal rate <sup>b</sup> (mm/yr)
Garlock	G1	$3.2 \pm 0.7$	$5.8 \pm 0.7$
Garlock	G2	$4.0 \pm 0.7$	$3.1 \pm 0.7$
Garlock	G3	$3.5 \pm 0.6$	$1.6 \pm 0.6$
Garlock	G4	$1.1 \pm 1.3$	$4.0 \pm 1.2$
Owens Valley	OV1	$-5.0 \pm 0.6$	$-1.9 \pm 0.7$
Owens Valley	OV2	$-4.6 \pm 0.5$	$-2.7 \pm 0.6$
Airport Lake	AL	$-5.3 \pm 0.7$	$-0.3 \pm 0.7$
Panamint Valley	PV	$-2.5 \pm 0.8$	$-1.9 \pm 1.0$
Hunter Mountain	HM	$-3.2 \pm 1.0$	$0.1 \pm 0.9$
Saline Valley	SV	$-2.7 \pm 0.9$	$-1.7 \pm 1.0$
Eureka Valley	EV	$-1.1 \pm 0.8$	$-3.0 \pm 0.9$
Furnace Creek	FC	$-2.7 \pm 0.6$	$0.8 \pm 0.6$
Fish Lake	FL	$-5.8 \pm 0.4$	$-0.4 \pm 0.4$
Death Valley	DV	$-2.8 \pm 0.5$	$-0.7 \pm 0.6$
Avawatz Mountains	AM	$-2.8 \pm 0.5$	$0.5 \pm 0.6$
Calico-Blackwater	CB1	$-1.7 \pm 1.5$	$-0.3 \pm 1.5$
Calico-Blackwater	CB2	$-1.4 \pm 1.4$	$-4.1 \pm 1.5$
Goldstone Lake	GL	$-6.4 \pm 1.7$	$1.3 \pm 1.7$
Eastern Mojave	EM1	$-2.8 \pm 0.5$	$-0.6 \pm 0.6$
Eastern Mojave	EM2	$-6.5 \pm 1.6$	$0.8 \pm 1.7$

<sup>a</sup>: negative values indicate right lateral motion

<sup>b</sup>: negative values indicate opening

<sup>5</sup> Available via anonymous ftp from <ftp://kosmos.agu.org>, directory "append" (Username "anonymous," Password "guest"); subdirectories are arranged by paper number.



**Figure 2.** Residual velocity field with 95% confidence ellipses. Labels are abbreviations for fault names. Slip rate estimates for each fault are given in Table 1.

the reciprocal of the squares of their error estimates. The reference frame of the solution is realized by requiring that the weighted mean of the residual velocities of the 17 sites on the eastern most block south of 37.4°N is zero. In general, horizontal motions of a block are described by three parameters: two velocity components and a rotation rate. Faults bounding a rotating block have slip rates that vary along strike. The elastic deformation associated with these faults is calculated by dividing them into short segments (one km or less) and using locally determined slip rates. An inversion carried out with translations only, in order to decrease the number of degrees of freedom, showed apparent unmodeled rotation of the western Mojave block. For the model presented here, only that block is permitted to rotate, with an inferred rate of  $2.1 \pm 0.8^\circ/\text{My}$  clockwise. Fault slip rate estimates are given in Table 1. The residual (observed – predicted) velocity field is shown in Figure 2.

## Discussion

The  $\chi^2$  per degree of freedom (dof) is 0.8, with a mean residual magnitude of 0.6 mm/yr for each velocity component. It appears that the data and the model are consistent and that we have used an appropriate number of free parameters.

Because unmodeled effects such as variations in elastic moduli can have effects on velocity gradients comparable to those of varying  $D$  (e.g., Hager *et al.*, 1999), we choose not to fine-tune the locking depths of individual fault segments; rather we find an optimal regional locking depth,  $D = 15$  km, and investigate the effects of varying it. The main effect of a

larger  $D$  is to put  $\sim 1$  mm/yr more slip on the faults bounding the Sierra Nevada, while a smaller  $D$  yields less slip on these faults. Variations of slip rates with  $D = 5$  to 25 km are at about the 1- $\sigma$  level about the slip rate estimates with  $D = 15$  km.

Our slip rate estimates for strike-slip motions on the faults north of the Garlock indicate that almost half of the ECSZ shear is concentrated on the Airport Lake/Owens Valley faults ( $5.3 \pm 0.7$ ,  $5.0 \pm 0.6$ , and  $4.6 \pm 0.5$  mm/yr on AL, OV1, and OV2), with  $2.5 \pm 0.8$  mm/yr on the Panamint Valley fault and  $2.8 \pm 0.5$  mm/yr on the Death Valley fault. Bennett *et al.* [1997] gave a comparable total for Death Valley plus Panamint Valley, but had insufficient data to resolve either their individual contributions or the rate of slip on the Owens Valley fault. Our rates are consistent, within the errors, with those of Gan *et al.* [2000], although they inferred a somewhat higher rate,  $6.9 \pm 1.6$  mm/yr, for the Owens Valley fault, probably because they ignored the shear associated with the SAF. Dixon *et al.* [2000] used a simple 2-D viscoelastic model to infer  $3 \pm 2$  mm/yr motion on OV1, also indistinguishable statistically. Beanland and Clark [1994] estimated a Holocene rate on the Owens Valley fault in the OV1-OV2 region of  $2 \pm 1$  mm/yr by assuming that the 3 most recent earthquakes there had the same average slip that they found for the 1872 event, 6 m. If not all the displacement on basement faults penetrates through the thick sediment cover to the surface, the 10 m maximum surface displacement might provide a better estimate of basement offset, implying a rate of  $3 \pm 2$  mm/yr, consistent with the geodetic rate.

We estimate  $5.8 \pm 0.4$  mm/yr slip on the Fish Lake Valley (FL) fault zone, consistent with the  $8 \pm 2$  mm/yr of Dixon *et al.* [2000]. The Quaternary rate on the FL segment is 6-12 mm/yr, with evidence for deceleration to 1-3 mm/yr over the past 50-75 ka, with slip transferred to faults to the east via the Emigrant Peak (EP) fault zone [Reheis and Sawyer, 1997]. The three sites northeast of FL exhibit coherent residual velocities. Because we have only one site to the NW of EP, we do not invert for the motion of a separate block, but the residual velocities of the sites north of FL suggest  $\sim 3$  mm/yr of slip is transferred to EP, which would reduce the slip on FL.

We also estimate the fault-perpendicular component of motion. The divergence between the blocks increases from east to west, from  $0.7 \pm 0.6$  mm/yr on block boundary DV to  $1.9 \pm 1.0$  mm/yr on PV, to  $2.7 \pm 0.6$  mm/yr on OV2. Divergence should promote volcanism and geothermal activity. The rapid divergence on OV1 and OV2 may help to explain the high rate of volcanic activity in the Owens Valley region and geothermal activity to the south.

The conclusion that the Owens Valley/Airport Lake faults have higher rates of slip than the Death Valley fault differs substantially from that of Hearn and Humphreys [1998], who inferred just the opposite. However, they had only the velocity of a single VLBI site, Mojave, to constrain the distribution of slip in this region.

South of the Garlock fault, the locations of boundaries CB2, EM1, and EM2 are poorly constrained, although the surface faulting, focal mechanisms, and aftershock distributions of the Landers and Hector Mine earthquakes suggest the appropriateness of our choice of boundaries. To avoid transient velocities associated with the Landers earthquake, we do not include GPS data in this region. Our model predicts  $\sim 11$  mm/yr of right lateral motion across the southern Mojave at latitude 34.5°, consistent with previous geodetic studies of the southern Mojave based on triangulation [Savage *et al.*, 1990; Sauber *et al.*, 1994],

demonstrating the utility of the path integral constraint implicit in the block model. The total right lateral slip budget in the southern Mojave in our model is consistent with the preferred 6 - 10 Ma geologic rate of 6-12 mm/yr given by *Dokka and Travis* [1990].

West of its intersection with the Calico-Blackwater fault, the Garlock has  $3.5 \pm 0.6$  mm/yr of slip. *Miller et al.* [2001], using fewer data and a dislocation model that does not include block rotation or satisfy path integral constraints, concluded that the slip rate on the western Garlock is only 1-2 mm/yr. Our estimate is consistent with the Holocene rate of 4 - 7 mm/yr [*McGill and Sieh*, 1991]. Both are less than the paleoseismologic estimate of 10-12 mm/yr, perhaps resulting from a cluster of earthquakes over the past 1.5 ka [*McGill and Rockwell*, 1998]. To the east of this intersection we estimate the left lateral strike slip rate on G4 to be  $1.1 \pm 1.3$  mm/yr, significantly less than the Holocene estimate of 4 - 9 mm/yr [*McGill and Sieh*, 1993]. The residual velocities of the 3 sites on the block to the south of G4 suggest that unmodeled rotation could explain this discrepancy.

*Ron et al.* [2001] proposed a model in which the western Mojave undergoes simple shear on faults parallel to the SAF, with negligible rotation. We tested this model by splitting the western Mojave block in two and allowing these two blocks to translate without rotating. Despite having one additional degree of freedom, this model has a  $\chi^2$  for the 9 sites in the western Mojave 19% greater than the model with a single rotating western Mojave block. In addition, the inferred slip rate of the western Garlock is  $2 \pm 2$  mm/yr, in worse agreement with the geologic rate. We cannot reject the *Ron et al.* model, although we prefer the rotating block model because it fits the data better with one less parameter.

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B. H. Hager, R. W. King, S. M. McClusky, B. J. Meade, Rm 54-614, Department of EAPS, Massachusetts Institute of Technology, Cambridge, MA 02139. (e-mail simon@chandler.mit.edu)  
 S. Bjornstad, F. Monastero, Geothermal Program Office, Code 83G000D, NAWA, China Lake, CA 93555  
 M. M. Miller, Dept. of Geological Sciences, Central Washington University, Ellensburg, WA 98926  
 B. J. Souter, Souter Computing Services, Baltimore, MD, 21239.

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